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Geostatistical interpolation of AIRS mid-tropospheric CO2

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- Introduction
- Fixed Rank Kriging (FRK)
- 9-year AIRS CO2 movies
- Conclusions



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Introduction

• Motivation: spatial interpolation of remote sensing data.



moving window average, 5 degree tophat

- Spatial interpolation is a necessary fact of life in analyzing remote sensing CO2.
- Examples include simple "gridding", linear interpolation, and moving window averaging, etc.
- Not all spatial interpolation methods provide measures of uncertainty. A spatial statistical model makes this possible.



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Let Z be an N-dimensional vector of observed AIRS values at locations (s_1, s_2, \dots, s_N) .

We estimate the AIRS value at location s_0 with the following form,

$$\hat{\mathbf{Y}}(\mathbf{s}_0) = \mathbf{a}(\mathbf{s}_0)' \mathbf{Z},$$

where $\mathbf{a}(s_0)$ is an N-dimensional vector of kriging coefficients. We wish to find the vector $\mathbf{a}(s_0)$ that minimizes the expected squared error,

 $E([Y(s) - a(s_0)' Z]^2).$



Jet Propulsion Laboratory California Institute of Technology Pasadena, California **Traditional Kriging**

Solving for the kriging coefficient vector using matrix derivatives, we get,

 $\mathbf{a}(s_0) = \mathbf{\Sigma}^{-1} \mathbf{c}(s_0)',$

where

 $\Sigma^{-1} = Var(\mathbf{Z})$ is an NxN matrix, $\mathbf{c}(s_0)' = Cov(\mathbf{Z}, Y(s_0))$ is an Nx1 vector.

• To compute the kriging coefficients, $\mathbf{a}(s)$, we need to compute the inverse of the $(N \times N)$ covariance matrix $\boldsymbol{\Sigma} = var(\boldsymbol{Z})$.



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Fixed ranked kriging (FRK) models the covariance structure with the following form,

 $cov(Y(s_1), Y(s_2)) = S(s_1)' K S(s_2),$

which leads to the following form for the covariance matrix Σ ,

 $\Sigma = S' K S + D,$

Because of this special form, Σ can be inverted quickly using the Sherman-Morrison-Woodbury formula,

 $\Sigma^{-1} = D^{-1} - D^{-1} S' (K^{-1} + S D^{-1} S')^{-1} S D^{-1}.$

- **S**(s) is an r-dimensional spatial basis expansion of s,
- **K** is an (r x r) matrix,
- S is an $(r \times N)$ matrix of $S(\bullet)$ evaluated at all observation location,
- **D** is an $(N \times N)$ diagonal matrix of measurement-error variance.



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FRK can handle massive amount of data with low computational burden.

- FRK properly accounts for the spatial dependence between observations,
- It produces estimates of prediction error, which allow for hypothesis testing,
- It is well-suited for producing Level 3 data,
- In this presentation, we apply FRK to global AIRS CO2 record (September 2002 January 2012), using 9-day moving window.



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FRK Outputs









Sample FRK input and outputs for 1 time period: November 21, 2010.

Top left:raw input dataTop right:FRK CO2 estimatesBottom left:FRK error estimates



FRK movies

Sep 05, 2002





Sep 05, 2002



Top left:	AIRS	raw	CO2 ma	ар
Top right:	AIRS	FRK	map	
Bottom left:	AIRS	FRK	error	map



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Some notes:

- The higher uncertainties in the Bay of Bengal, Africa, and South America are due to lower yields.
- The increased uncertainty in the northern latitudes in 2011 may be due to loss of L2 yield as a result of the steady degradation of AMSU channel 5.
- The belt of elevated CO2 in the Southern Hemisphere is an annually recurring feature.
- This cylindrical projection distorts the Arctic region, leading to an exaggerated perception of data sparseness in the region due to the fact that the resolution of the output is $1^{\circ} \times 1^{\circ}$.



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- Smoothed maps of AIRS CO2 does well in highlighting movement and evolution of mid-tropospheric CO2 throughout the 9-year timespan.
- FRK can provide a new type of Level 3 product with the following properties:
 - No spatial gap and a temporal resolution of a few days.
 - Outputs may be customized to match any desired spatial resolution.
 - The technique may be generalized to work in the third dimension (altitude) and employed for other physical parameters.



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Appendix A: projection distortion









Nov 21, 2010

