Observational evidence of frontal-scale atmospheric responses to Kuroshio Extension variability

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Ocean has slow and frontal scale processes



Semi-permanent meanders.

For the Kuroshio Extension (KE), interannual-to-decadal variability is pronounced in terms of its location and strength.

What is the imprint of KE variations on local atmosphere at low frequencies?

Kuroshio Extension (KE) index

- SSHA from AVISO averaged over the domain, 31-36°N and 140-165°E.
- Apply 13-month moving average.
- DJF of 1993-2012 (Dec in 1993-2012 and Jan-Feb in 1994-2013).
- Positive: strengthen eastward transport and migrate northward.
- Decadal variability.

• Normalized KE index



Data and methods used

- Data: monthly-mean in DJF
- 1) Satellite
 - TRMM 3B43 (50°S-50°N) surface precipitation, TRMM 3A12 (40°S-40°S) rain rate and latent heating rate (December 1998 to February 2013)
 - AIRS level 3 air temperature and geopotential height (December 2003 to February 2013)
- 2) Reanalysis
 - ERA-Interim: 10 m wind vectors, horizontal and vertical velocity, sea level pressure, air temperature and geopotential height
- Methods
 - Spatial filter: a running window with 10°lon x 2°lat
 - ENSO signal removal
 - Linear regression of the KE index onto climate variables pointwise

Manifestation of KE index on SST variability

• Basin scale



Manifestation of KE index on SST variability

Frontal scale: warm SST anomaly – SST ridges; cold SST anomaly – SST troughs.



Atmospheric response



Convergence of 10m wind w/ mean 10m wind vectors



Displacement may be linked to the advection effect of background flows.

Pressure adjustment mechanism: KE-induced SST-> pressure-> surface convergence-> vertical motion-> precipitation

Atmospheric response



Atmospheric response Vertical profiles along 35°N



0.5

0.4 0.3

0.2 0.1

0

-0.1

-0.2

-0.3

-0.4 -0.5

Atmospheric response Vertical profiles along 35.5°N



Atmospheric response Vertical profiles along 35.5°N



- Consistent imprints of the KE variations on the local atmosphere.
- The effect of the KE variations penetrates towards the free atmosphere is observed by the TRMM and AIRS data, which implies a potential pathway to influence the large-scale atmospheric circulation.

Atmospheric response Vertical profiles along 35.5°N



Compare AIRS and ERA-Interim air temperature



^{0.2} 0.16 • AIRS:

- The observing conditions in the absence of optically thick clouds.
- Incomplete diurnal cycle.
- ERAi:
 - Fewer constraints in the upper troposphere.
 - Responses are subject to prescribed SSTs.

Modeling bias in the vertical penetrating effect of small-scale ocean processes?



45[°] N

30[°] N

400 hPa

150[°] E

120[°] E

0

180[°] E

-0.05

Mean for DJF of 1993-2012 Shading: vertical velocity (Pa/s) Contours: geopotential height (m) Wind vectors



Horizontal convergence (1/s), positive: convergence



• The mean state of vertical velocity appears to act like a cap.

0.05

-0.05

0



Horizontal convergence (1/s), positive: convergence



Mean-state vertical velocity along 35N Mean-state vertical velocity along 35N Mean-state vertical velocity along 35N

- The mean state of vertical velocity appears to act like a cap.
- To balance air motion from the surface, air motion changes in the middle- and upper- troposphere.







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- To balance air motion from the surface, air motion changes in the middle- and upper- troposphere.

Summary

- Consistent local atmospheric responses to KE variations at low frequencies are found.
- Vertical penetrating effects of KE variations are present in the satellitebased observations.
- Discrepancies between the AIRS and ERA-Interim data suggest the imperfect interpretation of atmospheric responses to small-scale ocean processes in numerical models.
- Vertical tilted responses may be linked to the background flow effect.

Back-up slides



Approximate pressure converted from height





Atmospheric response



AIRS-derived geopotential height at 850hPa (K)