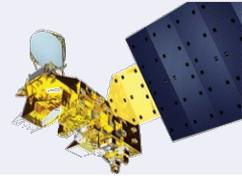


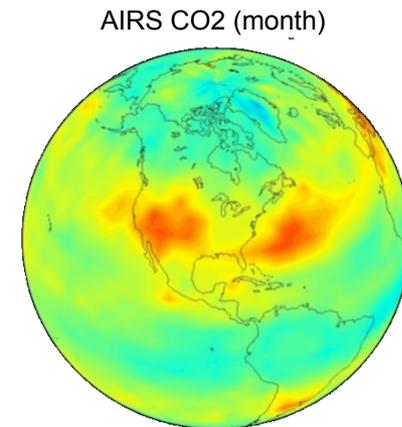
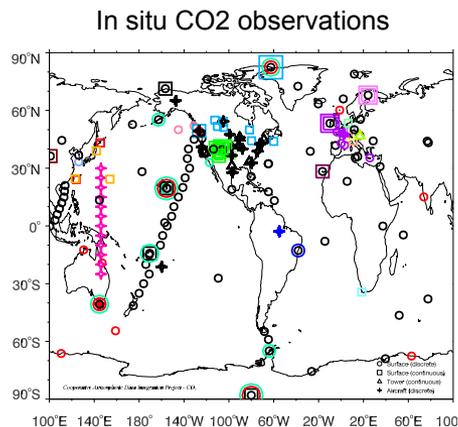
# AIRS CO<sub>2</sub> Science Highlights and Project Status

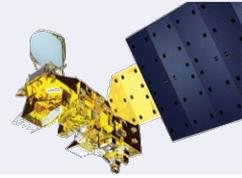
**Edward Olsen, Stephen Licata, Luke Chen,  
Thomas Pagano, Hai Nguyen, Alexander Ruzmaikin**  
California Institute of Technology, Jet Propulsion Laboratory  
**Xun Jiang**  
University of Houston, Houston Texas

**NASA Sounder Science Team Meeting  
17-19 March 2014**



- Why is the retrieval of free tropospheric CO<sub>2</sub> important?
  - Atmospheric CO<sub>2</sub> is the major agent for radiative forcing
    - Interannual increase due to human activity, and ~50% of anthropogenic CO<sub>2</sub> remains in atmosphere
  - The free troposphere is the pathway by which CO<sub>2</sub> produced at the surface by natural and anthropic processes is circulated around the globe to be deposited in the natural sinks
  - The free tropospheric (and stratospheric) CO<sub>2</sub> are the background which must be accounted for in the process of determining the surface carbon flux by remote sensing
    - SH, “the garbage dump for the NH CO<sub>2</sub> emissions”, very sparsely observed from ground/airborne
  - Modeling of atmospheric transport processes are yet not sufficiently developed
    - Vertical lofting
    - Inter-hemispherical transport
- Satellite retrievals provide the spatial and temporal coverage over the globe necessary to elucidate the transport of CO<sub>2</sub> and to ultimately identify the regional sources and sinks and net fluxes around the globe (AIRS, GOSAT, SCIAMACHY, TES, and forthcoming OCO-2)



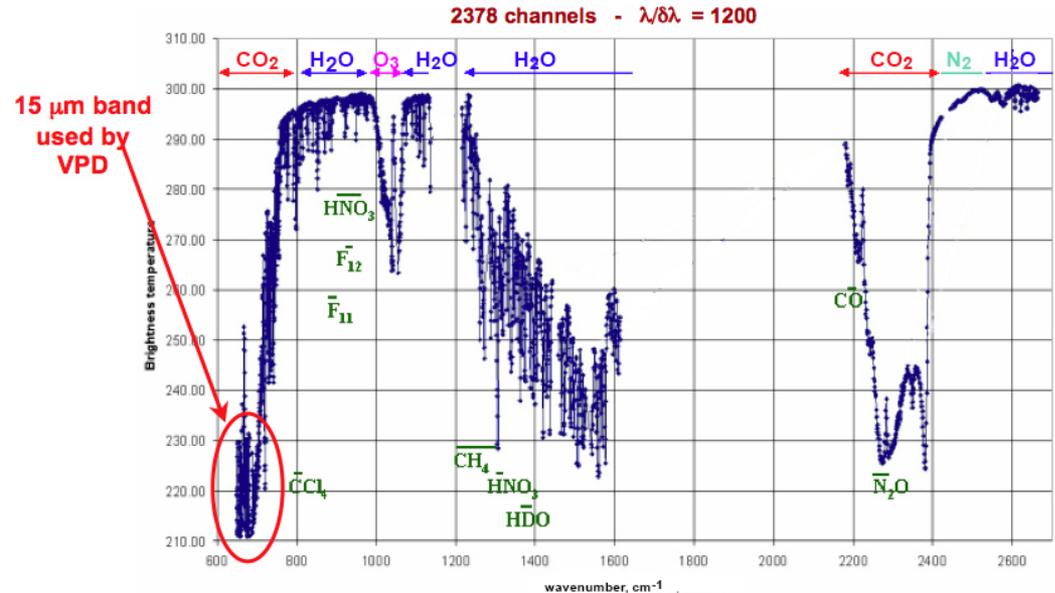


## AIRS Hyperspectral Infrared Spectrum

### AIRS Channels for Tropical Atmosphere with $T_{surf} = 301K$

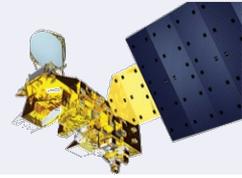
#### AIRS Characteristics

- **Aqua Spacecraft**
  - Launched: May 4, 2002
- **Sun-Synchronous A-train Orbit: 705 km**
  - 1:30 pm Equatorial ascending crossing
  - 1:30 am Equatorial descending crossing
  - 98.14° orbital inclination
  - Global Daily Coverage
- **2378 Channel Grating Spectrometer**
  - IR IFOV : 1.1° x 0.6°
  - 13.5 km spatial resolution at nadir
  - ±48.95° Cross-track scan
    - 90 IFOV per 2.667 sec
  - Spectral Range: 3.5-15.4 μm
  - L1B Radiances have Climate Quality Accuracy and Stability

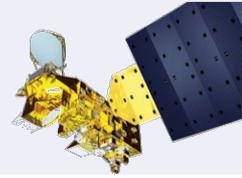


- AIRS  $CO_2$  retrieval algorithm employs method of Vanishing Partial Derivatives (VPD) ingesting AIRS L2 cloud cleared radiances and physical retrievals
  - Chahine, et al. (2005) "On the Determination of Atmospheric Minor Gases by the Method of Vanishing Partial Derivatives with Application to  $CO_2$ ", GRL, 32, L22803, doi:10.1029/2005GL024165.
- AIRS V5 Mid-Tropospheric  $CO_2$  data product from 9/2002-Present is available at:
  - <http://disc.sci.gsfc.nasa.gov/AIRS/data-holdings>

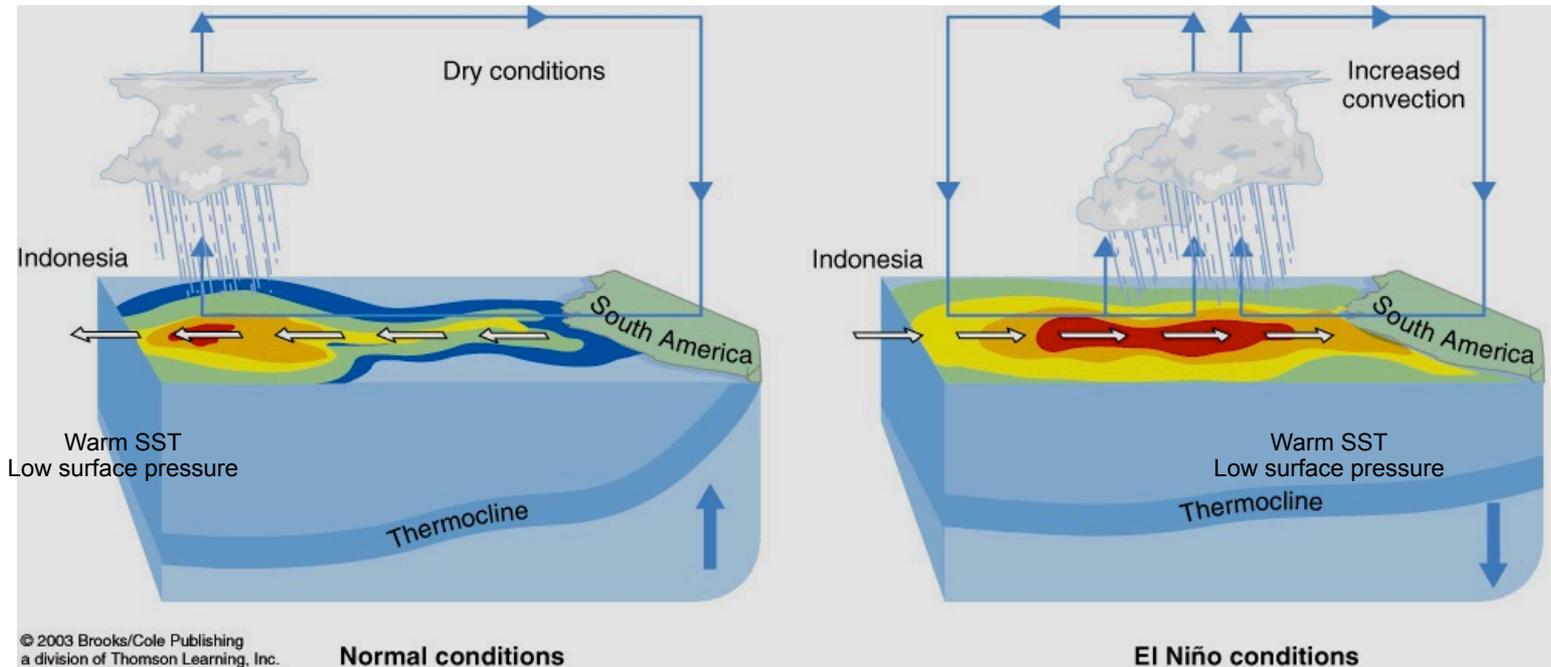
# Published Studies Bearing on CO<sub>2</sub> Transport Around the Globe



- Published studies demonstrate that the AIRS data contain signals arising from the large-scale circulation patterns in both the tropics and at high latitudes: ENSO, MJO, Walker Circulation, TBO, AO, and SSW in addition to the interannual growth of global CO<sub>2</sub> and its annual seasonal cycle.
  - Jiang, X et al. (2010), "Interannual variability of mid-tropospheric CO<sub>2</sub> from Atmospheric Infrared Sounder", *GRL*, 37, L13801, doi: 10.1029/2010GL042823
  - Li, KF et al. (2010), "Tropical mid-tropospheric CO<sub>2</sub> variability driven by the Madden–Julian oscillation." *Proceedings of the National Academy of Sciences* 107, no. 45, 19171-19175.
  - Wang, J et al. (2011), The influence of tropospheric biennial oscillation on mid-tropospheric CO<sub>2</sub>, *GRL*, 38, L20805, doi:10.1029/2011GL049288
  - Feng, et al. (2011) "Evaluating a 3-D transport model of atmospheric CO<sub>2</sub> using ground-based, aircraft, and space-borne data." *Atmospheric Chemistry and Physics* 11, no. 6, 2789-2803.
  - Liu, J et al. (2011), "CO<sub>2</sub> transport uncertainties from the uncertainties in meteorological fields." *Geophysical Research Letters* 38, no. 12.
  - Pagano, T et al. (2011) "Monthly representations of mid-tropospheric carbon dioxide from the Atmospheric Infrared Sounder." In *SPIE Optical Engineering+ Applications*, pp. 81580C-81580C. International Society for Optics and Photonics.
  - Kang, J-S et al. (2012), Estimation of surface carbon fluxes with an advanced data assimilation methodology, *J. Geophys. Res.*, 117, D24101, doi:10.1029/2012JD018259
  - Jiang, X et al. (2013) "Influence of El Niño on mid-tropospheric CO<sub>2</sub> from Atmospheric Infrared Sounder and Model." *Journal of the Atmospheric Sciences* 70, no. 1.
  - Jiang, X et al. (2013) "Influence of Stratospheric Sudden Warming on AIRS Mid-tropospheric CO<sub>2</sub>." *Journal of the Atmospheric Sciences* 70, no. 8.
  - Pagano, T, et al. (2014) "Global variability of mid-tropospheric CO<sub>2</sub> as measured by the Atmospheric Infrared Sounder", *Journal of Applied Remote Sensing* (accepted for publication).



ENSO is an important interannual variability in the tropical region and impacts the spatial/vertical distribution of tracers



El Niño condition - Eastward shift of warm sea surface temperature and low pressure and reversal of trade winds

Result – convection center also shifts eastward, thus there should be an impact upon the spatial distribution of concentration of CO<sub>2</sub> in the mid- to high troposphere.

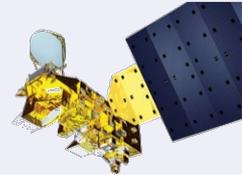
**Hypothesis: shift of large-scale vertical convection should be detected in AIRS CO<sub>2</sub>**



Jet Propulsion Laboratory  
California Institute of Technology

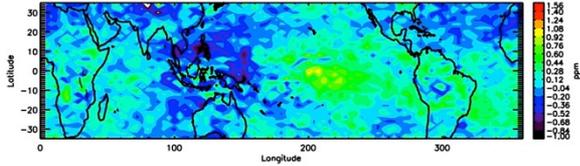
# El Niño Signal in AIRS Retrieved CO2

Jiang et al., (2012), "Influence of El Niño on mid-tropospheric CO2 from Atmospheric Infrared Sounder and model", JAS, doi:10.1175/JAS-D-11-0282.1

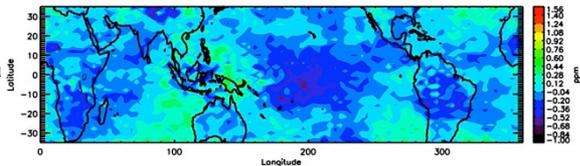


## AIRS CO2

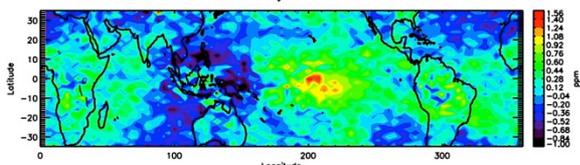
11 El Niño months



17 La Niña months

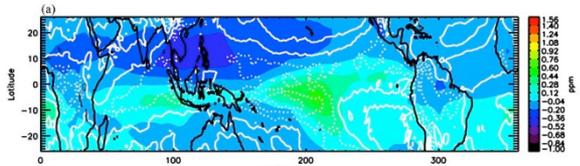


El Niño - La Niña

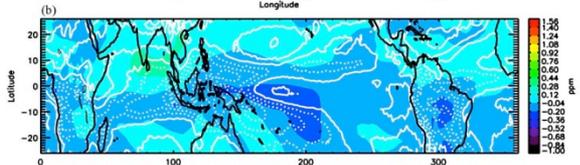


## MOZART-2 CO2

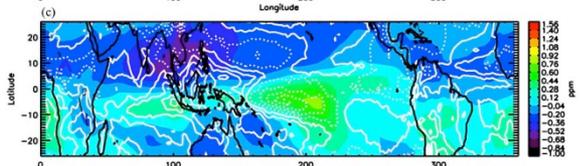
El Niño



La Niña

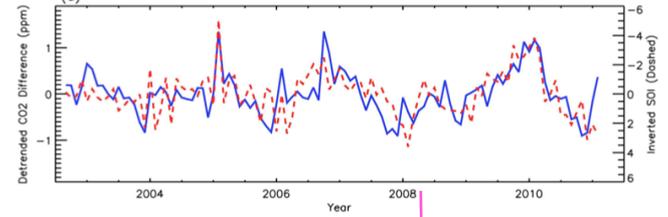


El Niño - La Niña

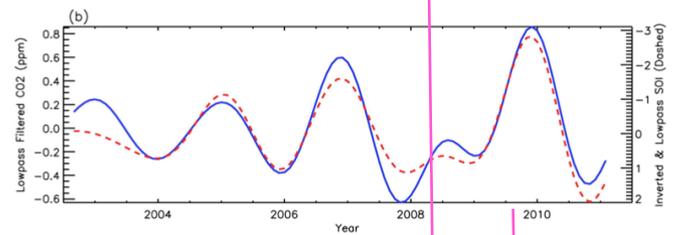


There is more (less) AIRS retrieved CO2 in the Central Pacific and less (more) AIRS retrieved CO2 in the Western Pacific during El Niño (La Niña) events.

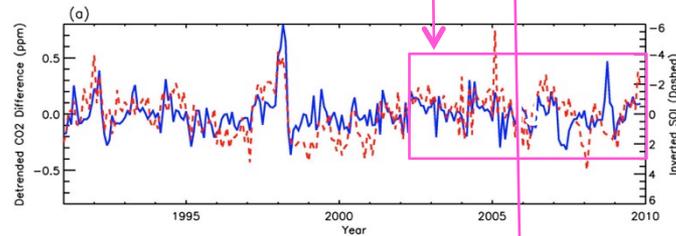
CO2 difference between the Central Pacific (190°E-240°E; 18°S-18°N) and the Western Pacific (110°E-160°E, 18°S-18°N) (Blue solid line)  
Inverted and detrended Southern Oscillation Index (Red dashed line)



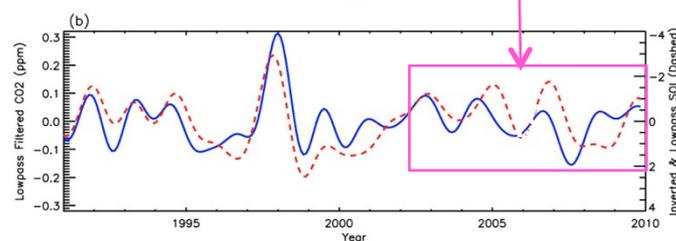
Raw Data  
Correlation:  
**0.62**



Low pass  
Correlation:  
**0.94**



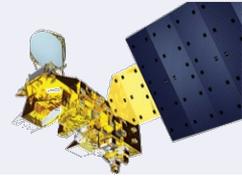
Raw Data  
Correlation:  
**0.48**



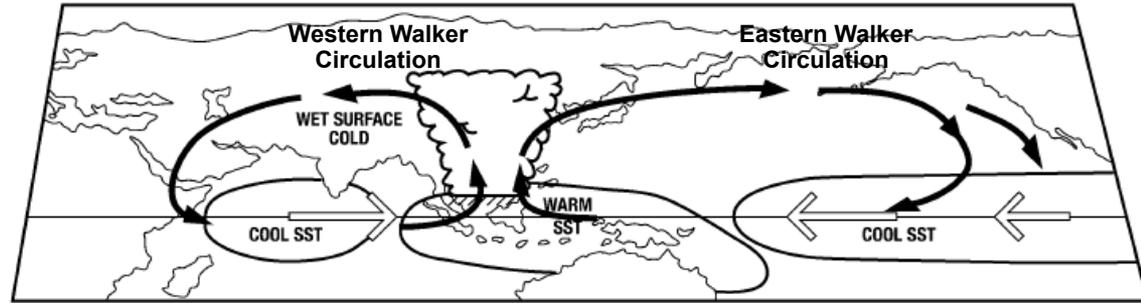
Low pass  
Correlation:  
**0.65**

Similar results are seen in MOZART-2 mid-tropospheric CO2, although the amplitude is reduced and spatial pattern less concentrated in the model compared to that in AIRS CO2 data.

# Tropospheric Biennial Oscillation (TBO)



The tropospheric biennial oscillation is the tendency for a relatively strong Asian monsoon to be followed by a relatively weak monsoon, and vice versa, with roughly a 2-3 year period. This variability in the tropical region also impacts the spatial/vertical distribution of tracers due to its effect on vertical transport.

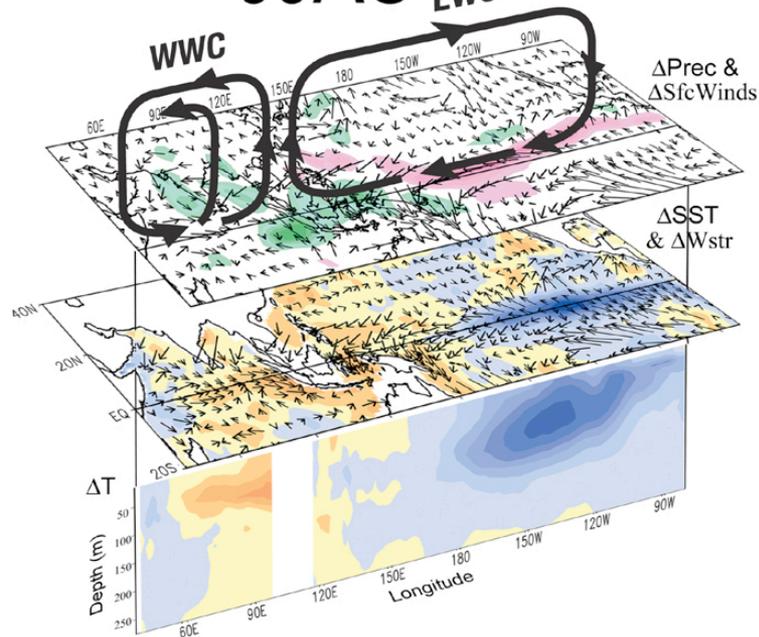


[Meehl & Arblaster, 2002]

JJAS

EWC

WWC



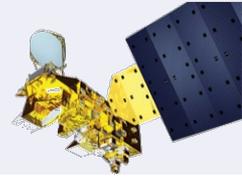
**Strong Monsoon - Strong Western Walker Circulation**

**Weak Monsoon - Weak Western Walker Circulation**

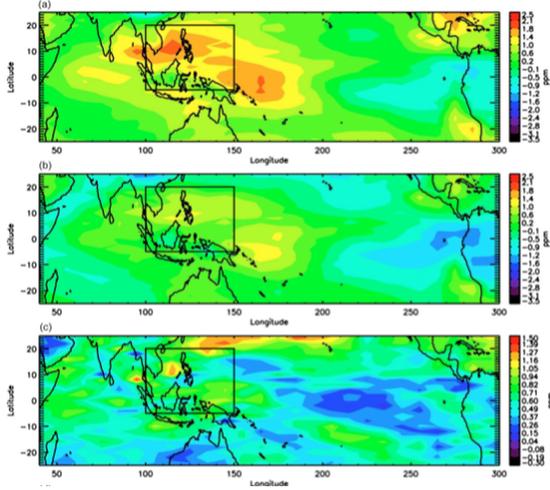
**Hypothesis: variable strength of Western Walker Center vertical convection should impact AIRS CO<sub>2</sub> product**



# TBO Signal in AIRS Retrieved CO2



## AIRS CO2



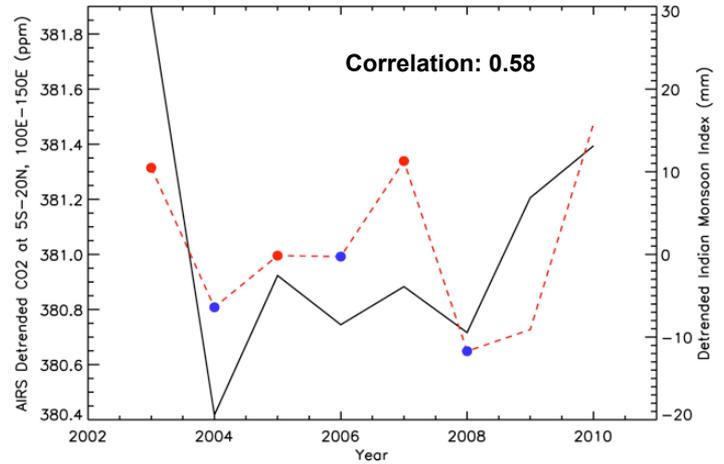
**Strong Monsoon (JJAS)**  
(2003, 2005, 2007, 2010)

**Weak Monsoon (JJAS)**  
(2004, 2006, 2008)

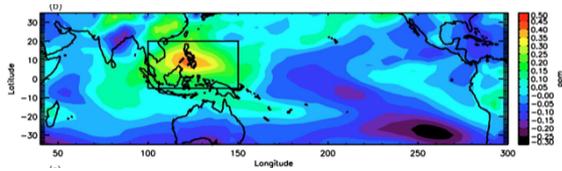
**Strong - Weak Monsoon**

There is more (less) AIRS retrieved CO2 in the Central Pacific and less (more) AIRS retrieved CO2 in the Western Pacific during weak (strong) Asian monsoon events.

**Black Line: AIRS detrended mid-trop CO2 at 5°S-20°N, 100°E-150°E**  
**Red Dashed Line: Detrended Indian Monsoon Index Derived from TRMM Precipitation (red points strong monsoon; blue points weak monsoon)**



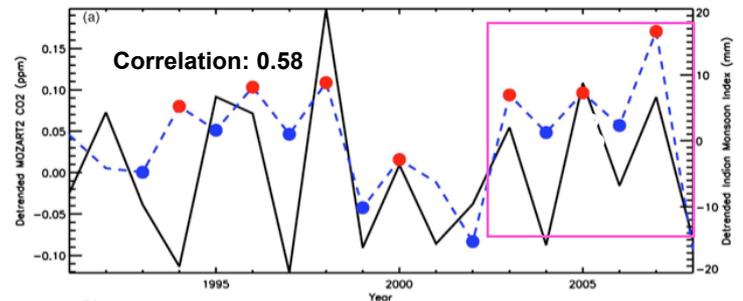
## MOZART-2 CO2



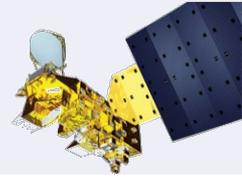
**Strong - Weak Monsoon**

Spatial distribution more restricted than AIRS concentrations. Again, more (less) CO2 in the Central Pacific and less (more) CO2 in the Western Pacific during weak (strong) Asian monsoon events.

**Black Line: MOZART-2 CO2 at 5°S-20°N, 100°E-150°E**  
**Blue Dashed Line: GPCP Monsoon Index (red points strong monsoon; blue points weak monsoon)**



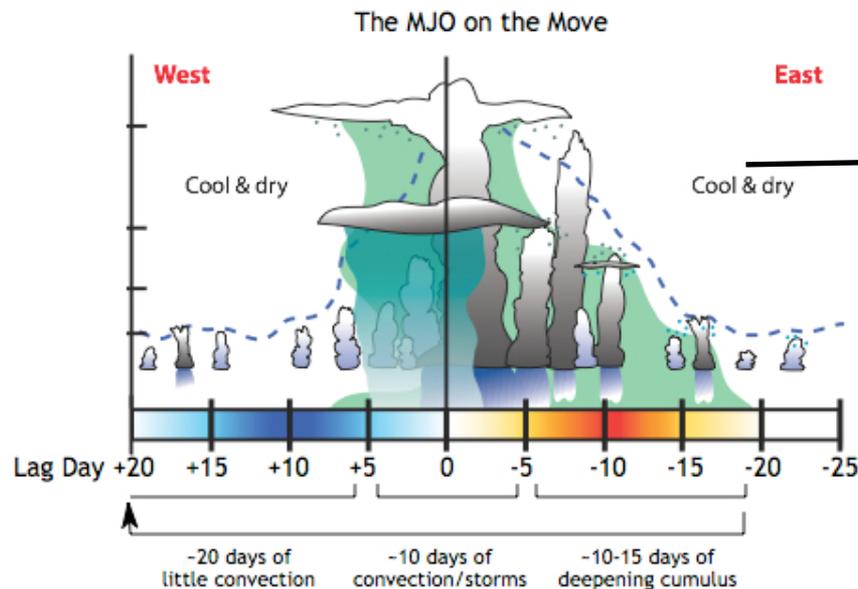
# Madden-Julian Oscillation (MJO)



The Madden–Julian oscillation (MJO) is the dominant component of intraseasonal variability in the tropics and has a period of 40 to 60 days. In the troposphere, the MJO is an equatorial planetary-scale wave envelope of multi-scale convective processes. Rather than being a standing pattern (like ENSO) it is a traveling pattern, propagating eastwards at approximately 4 to 8 m/s, through the atmosphere above the warm parts of the Indian and Pacific oceans. This overall circulation pattern manifests most clearly as anomalous rainfall.

The MJO is characterized by an eastward progression of large regions of both enhanced and suppressed tropical rainfall, observed moving eastward from the western Indian Ocean through the central tropical Pacific Ocean. This pattern of tropical rainfall then generally becomes nondescript as it moves over the cooler ocean waters of the eastern Pacific (except over the region of warmer water off the west coast of Central America) but sometimes reappears at low amplitude over the tropical Atlantic and higher amplitude over the Indian Ocean. The wet phase of enhanced convection and precipitation is followed by a dry phase where thunderstorm activity is suppressed.

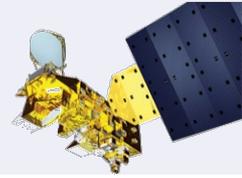
**Hypothesis: This enhanced large-scale convection and its movement eastward should create a signal visible in the AIRS retrieved mid-tropospheric CO<sub>2</sub> concentration.**



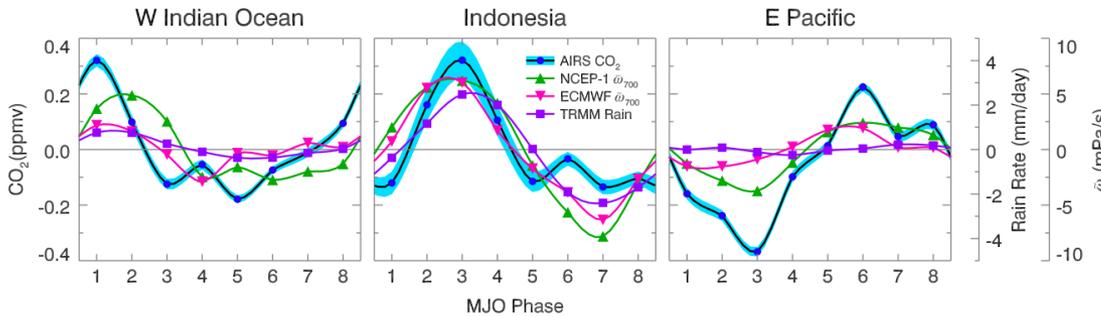
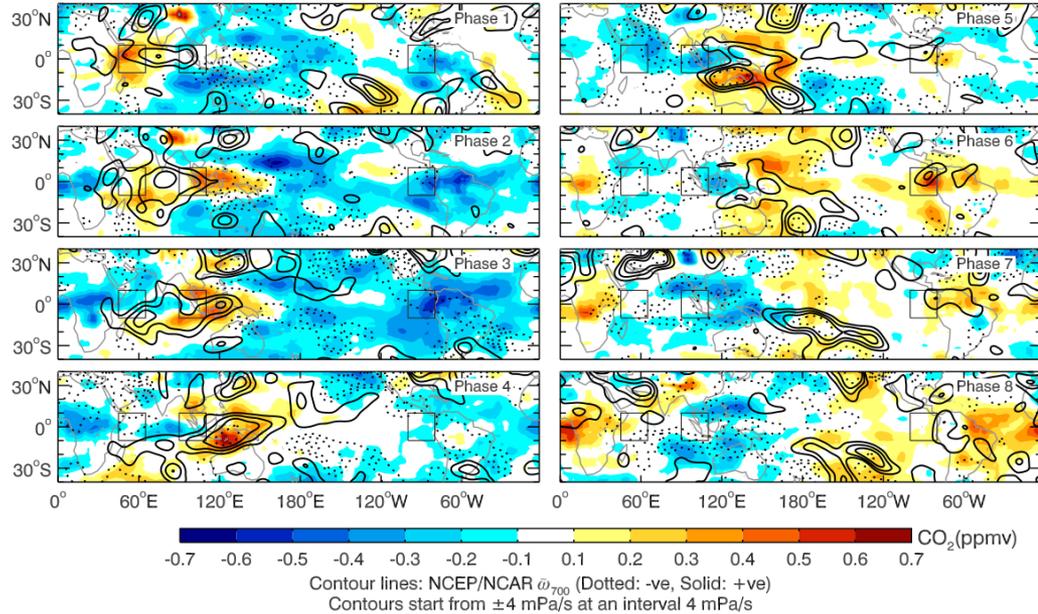


# MJO Signal in AIRS Retrieved CO<sub>2</sub>

Li et al. , (2010), "Tropical mid-tropospheric CO<sub>2</sub> variability driven by the Madden-Julian oscillation." *Proceedings of the National Academy of Sciences* 107, no. 45, 19171-19175.

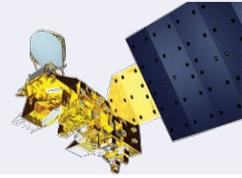


**MJO related AIRS CO<sub>2</sub> anomalies for 8 MJO 6-day phases. Data are for days with strong MJO activity,  $(RMM_1)^2 + (RMM_2)^2 \geq 1$ , for boreal winters in period from Nov 2002 through February 2010. Eastward propagation shown by solid/dotted MJO-related rainfall anomaly outside of threshold  $\pm 1$  mm/day.**



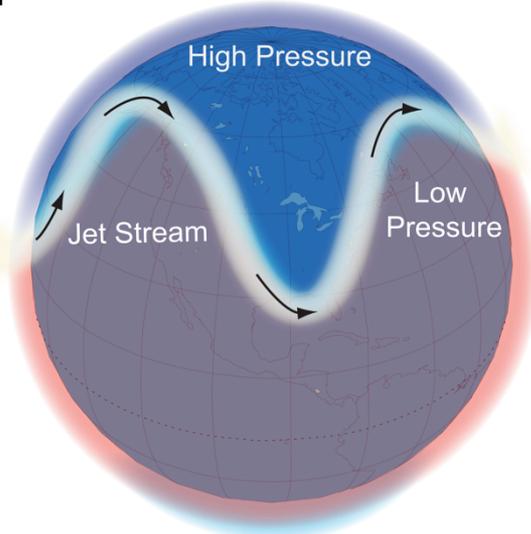
Composite MJO time series of AIRS CO<sub>2</sub>, reanalysis 700 hPa vertical velocity, and TRMM rainfall anomalies over the western Indian Ocean, Indonesia and the eastern Pacific Ocean enclosed by the square boxes in the upper figure. Over all three locations the CO<sub>2</sub> variation is positively correlated with ERA-interim  $\bar{\omega}_{700}$ . Positive CO<sub>2</sub> anomaly over regions characterized by rising air and negative CO<sub>2</sub> anomaly over regions characterized by sinking air appears to be generally true for the entire tropics.

# Northern Annular Mode/Arctic Oscillation (NAM/AO)

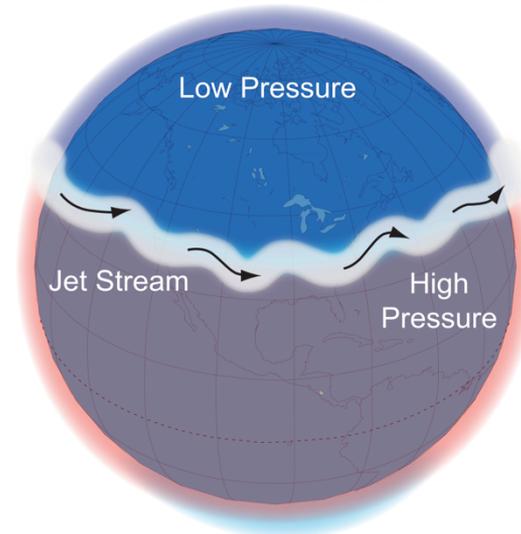


The Northern Annular Mode, or Arctic Oscillation, is characterized by pressure anomalies of one sign in the Arctic with the opposite anomalies centered about 37N-45N. These pressure patterns impact the middle latitude jet stream and thus the degree of isolation of the polar atmosphere from that of the middle latitudes and the height of the tropopause in the Arctic. Positive AO isolates the polar atmosphere and depresses the altitude of the tropopause.

Negative Phase



Positive Phase



**Negative AO – weak zonal winds, polar air mixes with mid-latitudes, higher altitude tropopause over the Arctic, frigid weather outbreaks in the mid-latitudes**

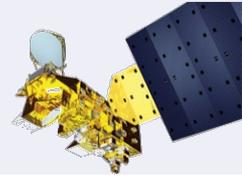
**Positive AO – strong zonal winds, isolated polar air, lower altitude Arctic tropopause**

**Hypothesis: Isolation of interior of polar vortex and downward intrusion of stratospheric air should decrease AIRS retrieved CO<sub>2</sub> over the Arctic.**



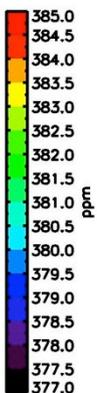
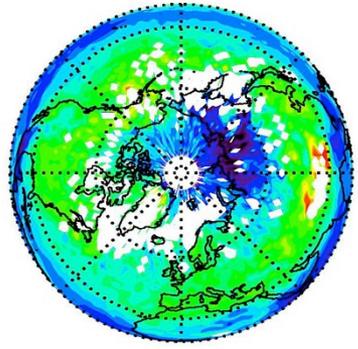
# Influence of NAM/AO on AIRS Retrieved CO2

Chang et al., (2010), "Interannual variability of mid-tropospheric CO2 from Atmospheric Infrared Sounder", GRL, 37, L13801, doi: 10.1029/2010GL042823 .

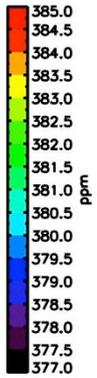
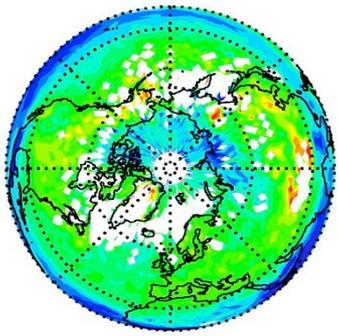


## AIRS CO2

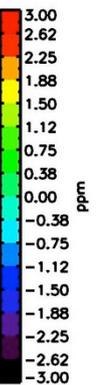
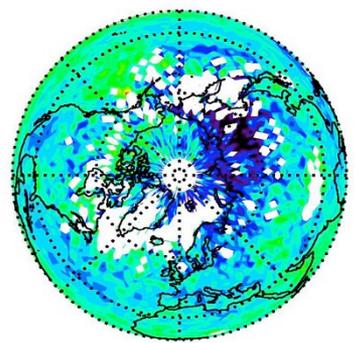
Strong Vortex  
-  
Positive AO Index (Nov-Apr)  
(2005, 2007)



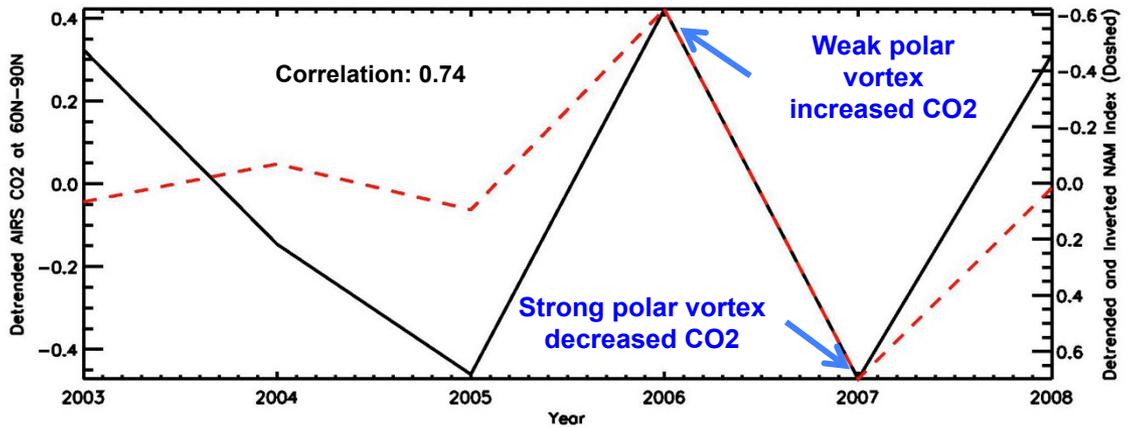
Weak Vortex  
-  
Negative AO Index (Nov-Apr)  
(2006, 2008)



Strong - Weak Vortex



Black Line: AIRS detrended mid-trop CO2 60N-90N (Nov-Apr)  
Red Dashed Line: Detrended and inverted Arctic Oscillation Index (Nov-Apr)

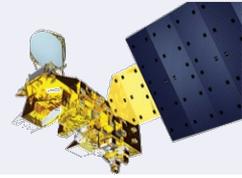


Strong vortex isolates the polar regions from the mid-latitudes and the tropopause sinks.

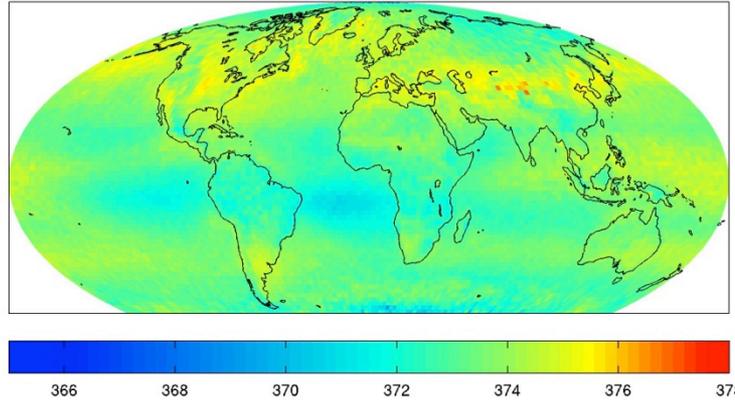
AIRS retrieved CO2 over the Arctic is decreased due to isolation from CO2-richer mid-latitudes and intrusion of lower-CO2 stratosphere.

*During a sudden stratospheric warming event, the process is rapidly reversed, and retrieved Arctic CO2 increases 2-3 ppm.*

# Persistent Depleted Region in of AIRS Mid-Tropospheric CO2 in South Atlantic Related to Eastern Walker Circulation

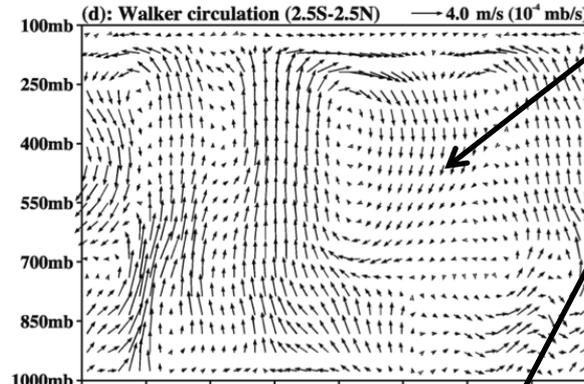
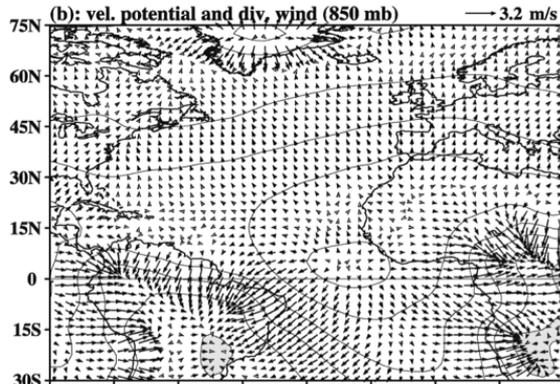


Annual Average  
2003-2010  
Detrended  
AIRS Mid-Trop CO2

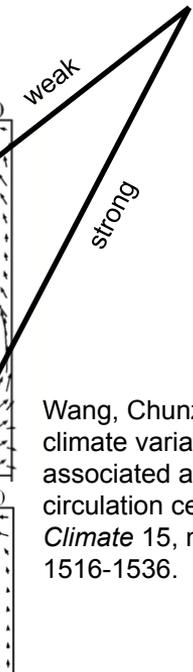
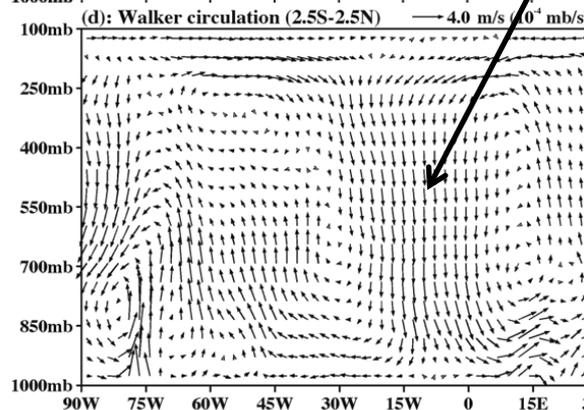
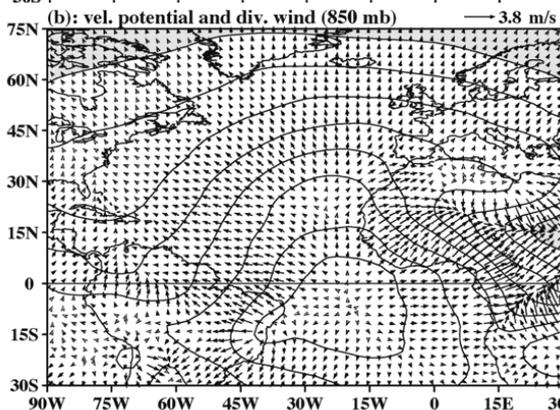


Downward Flow of  
Eastern Walker  
Circulation

January

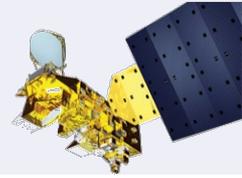


July



Wang, Chunzai. "Atlantic climate variability and its associated atmospheric circulation cells." *Journal of Climate* 15, no. 13 (2002): 1516-1536.

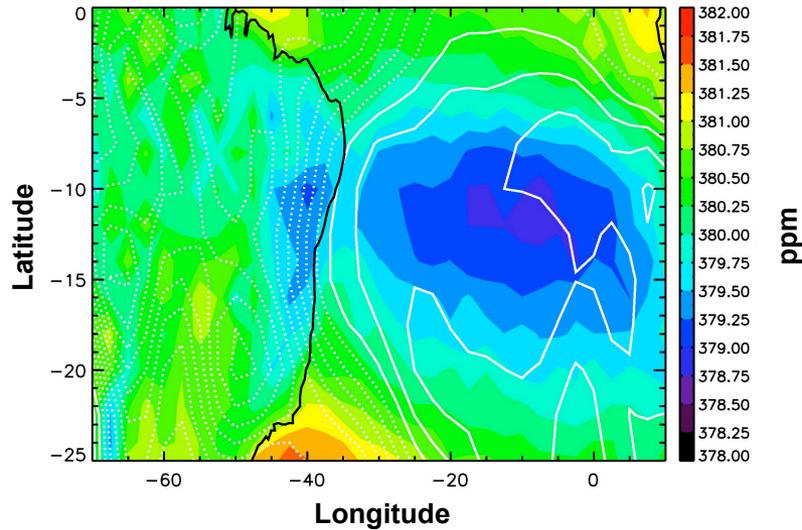
# Influence of Atlantic Walker Cell on AIRS Retrieved CO<sub>2</sub>



(Jiang et al. paper in preparation)

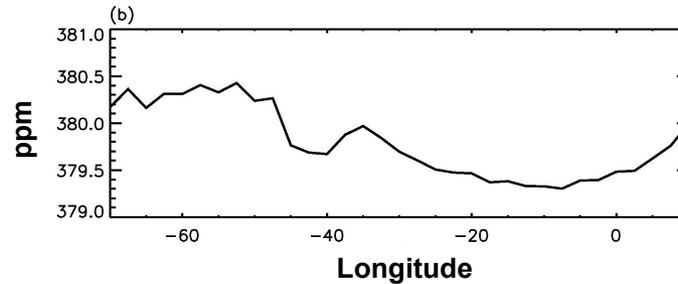
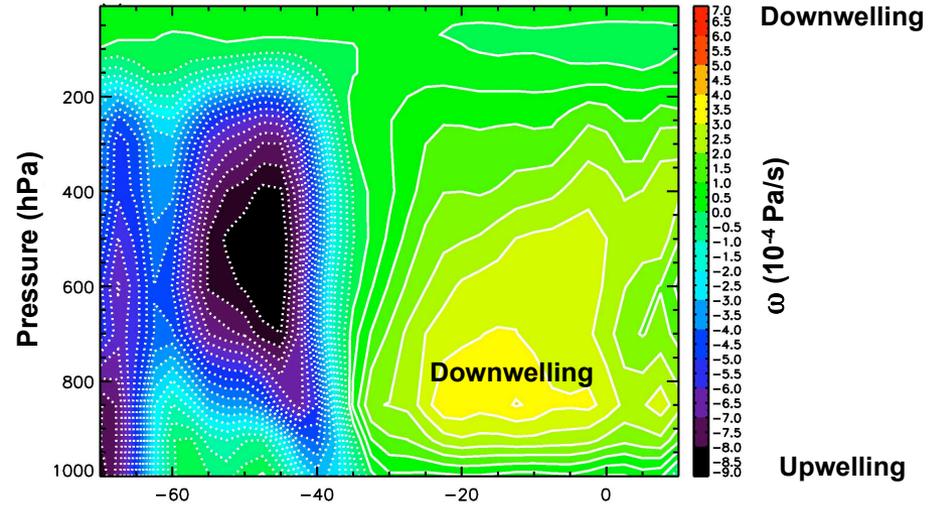
**Sinking air in the descending branch of the Atlantic Walker Cell brings lower concentration of CO<sub>2</sub> from the higher altitudes**

Average AIRS CO<sub>2</sub> in 2003-2010 DJFM



White contours are the NCEP2 400 hPa vertical pressure velocity ( $dP/dt$ ). Solid white contours refer to the sinking air. Dashed white contours refer to the rising air.

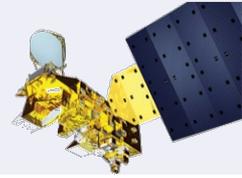
Omega ( $dP/dt$ ) in 2003-2010 DJFM Averaged over 20°S – 5°S



Variation with longitude of AIRS CO<sub>2</sub> averaged over 20°S-5°S in 2003-2010 DJFM.

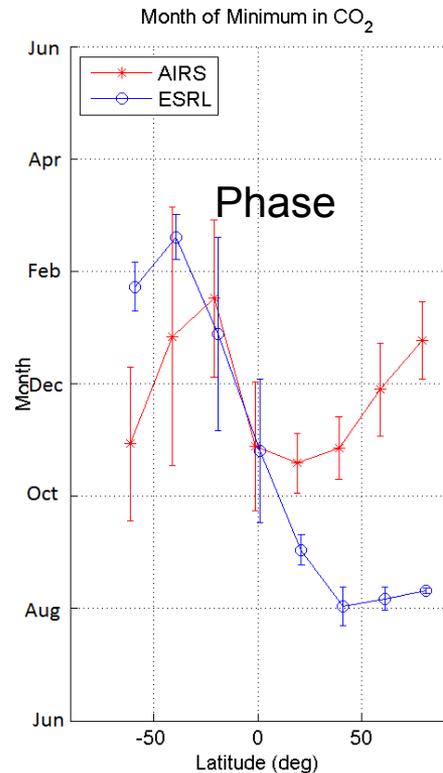
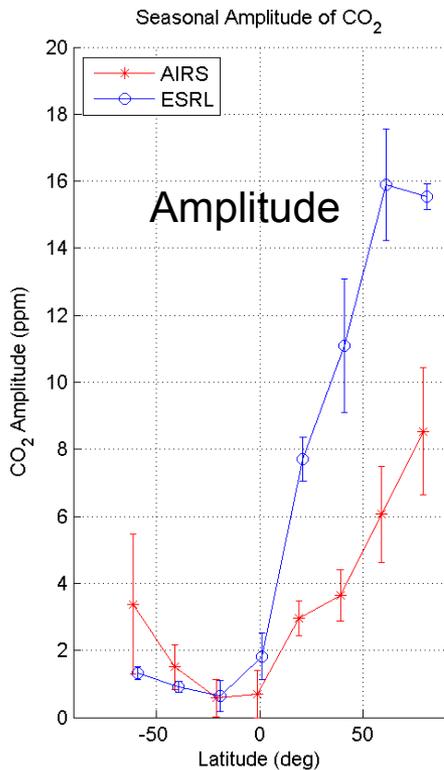
# AIRS CO<sub>2</sub> Shows Significant Influence of Surface CO<sub>2</sub> Seasonal Variability

(Pagano et al. (2014) "Global variability of mid-tropospheric CO<sub>2</sub> as measured by the Atmospheric Infrared Sounder", Journal of Applied Remote Sensing, (accepted).)

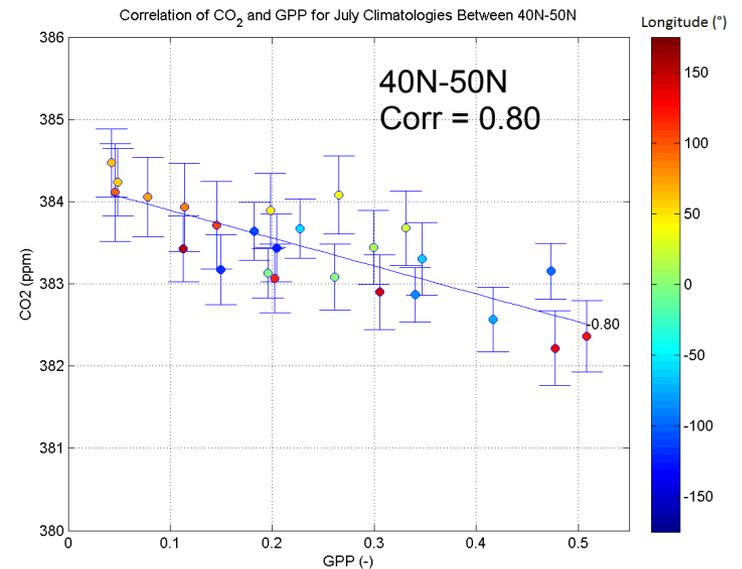
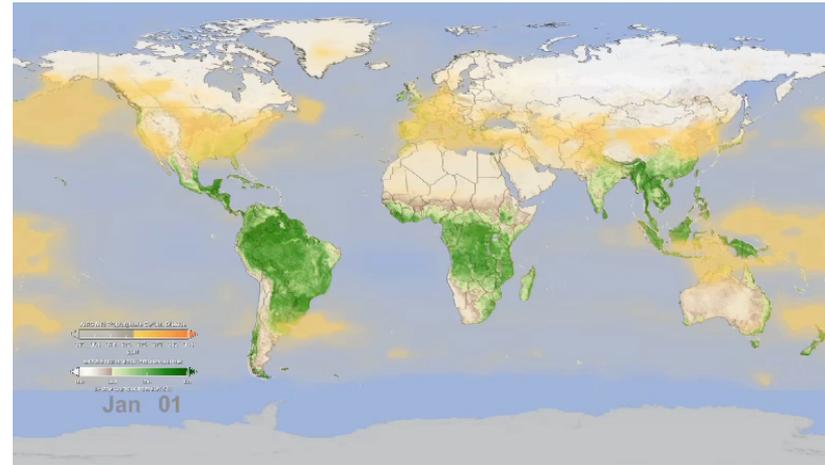


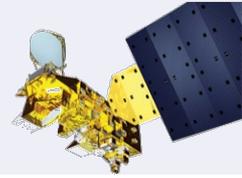
## AIRS CO<sub>2</sub> Seasonal Cycle:

- Strong seasonal cycle present in mid-trop
- NH damped and delayed compared to surface CO<sub>2</sub> concentration variation => vertical transport
- SH variation higher amplitude and precedes surface CO<sub>2</sub> variation => Inter-hemispherical transport



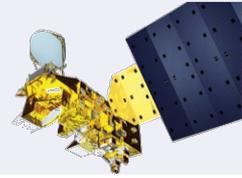
## High Correlation of CO<sub>2</sub> and GPP for July in NH Boreal Forests



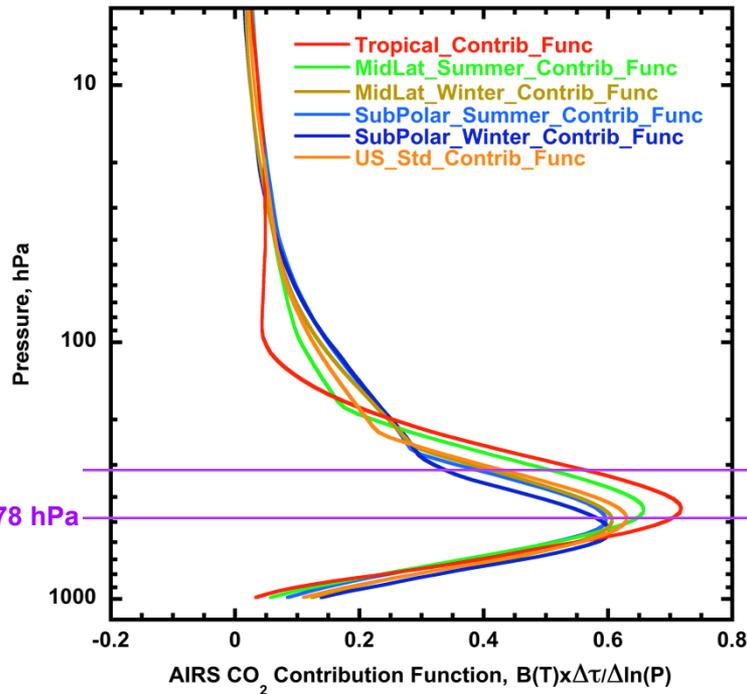


- V6 Prototype code
  - Code designed for rapid conversion to PGE
    - Execution options set via environmental variables
      - Channel lists, priors, SARTA version, QA filtering rules and thresholds
  - Capable of ingesting V5 and V6 AIRS L2 products
  - Capable of using three RTA versions (V107, V108 or V6)
  - Supports multiple channel sets for separate partial column retrievals
- V6 Prototype testing
  - V5 AIRS-Only CO<sub>2</sub> product released while V6 in development
  - Retrievals using same channel set consistent with V5 for  $|\text{lat}| < 50^\circ$ 
    - V6 does not require V5 radiance bias correction
    - Expanding and optimizing QA
    - Validating against in situ profiles

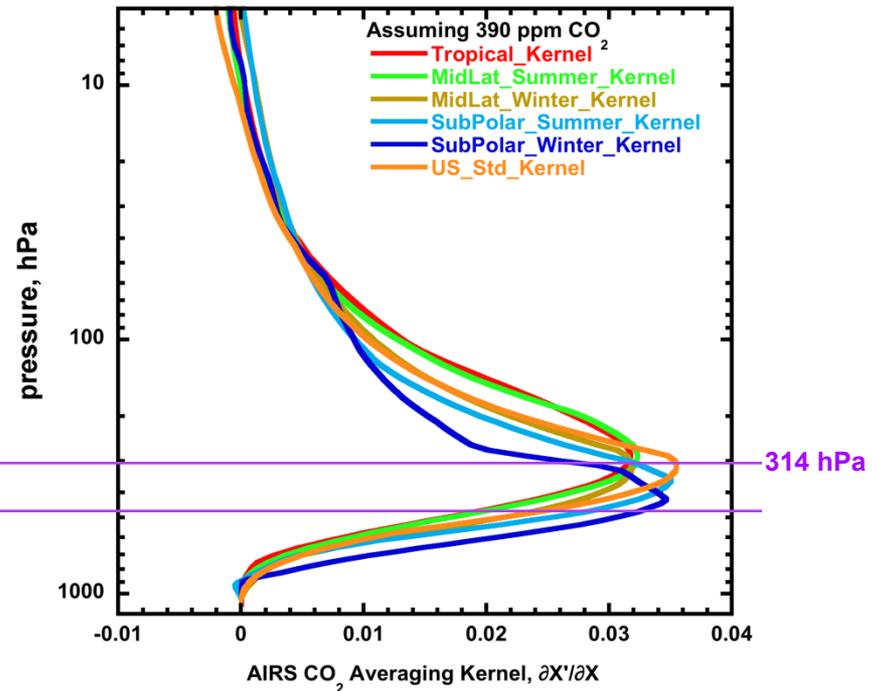
# AIRS Mid-Tropospheric CO<sub>2</sub> Product Contribution Function and Averaging Kernel



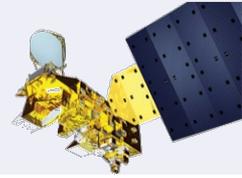
**Contribution Function**  
(used to choose V5 channels)



**Averaging Kernel**  
(used to assimilate product)

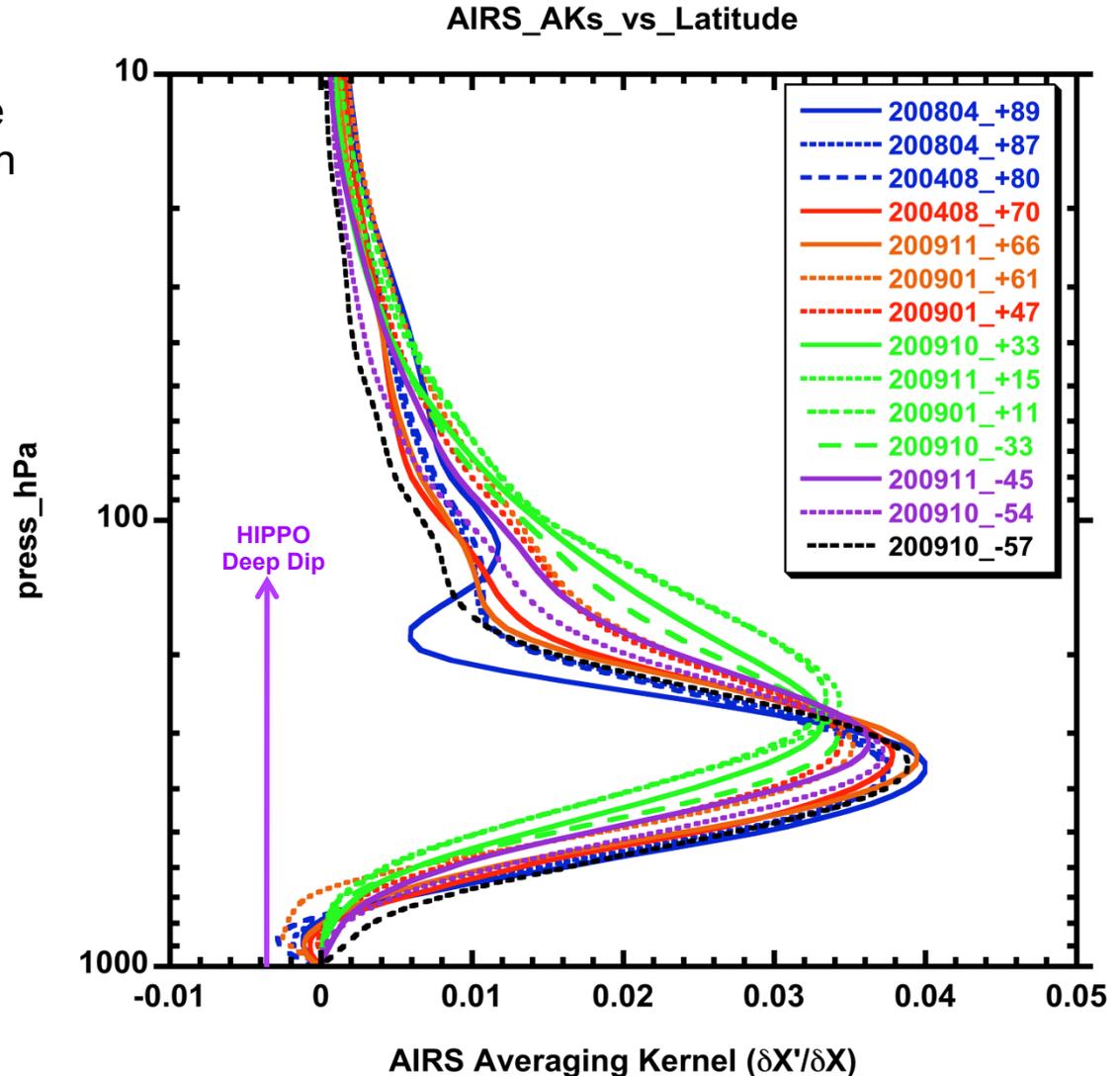


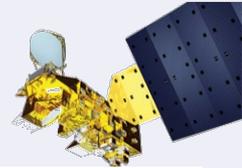
- The AIRS Mid-Trop Contribution Function is a measure of the contribution of an atmospheric layer to the TOA radiance used in the AIRS CO<sub>2</sub> retrieval
- The AIRS Mid-Trop Averaging Kernel is a measure of the sensitivity of the AIRS CO<sub>2</sub> retrieval to a change in CO<sub>2</sub> concentration in an atmospheric layer



Airborne profile measurements are the best available data for validation

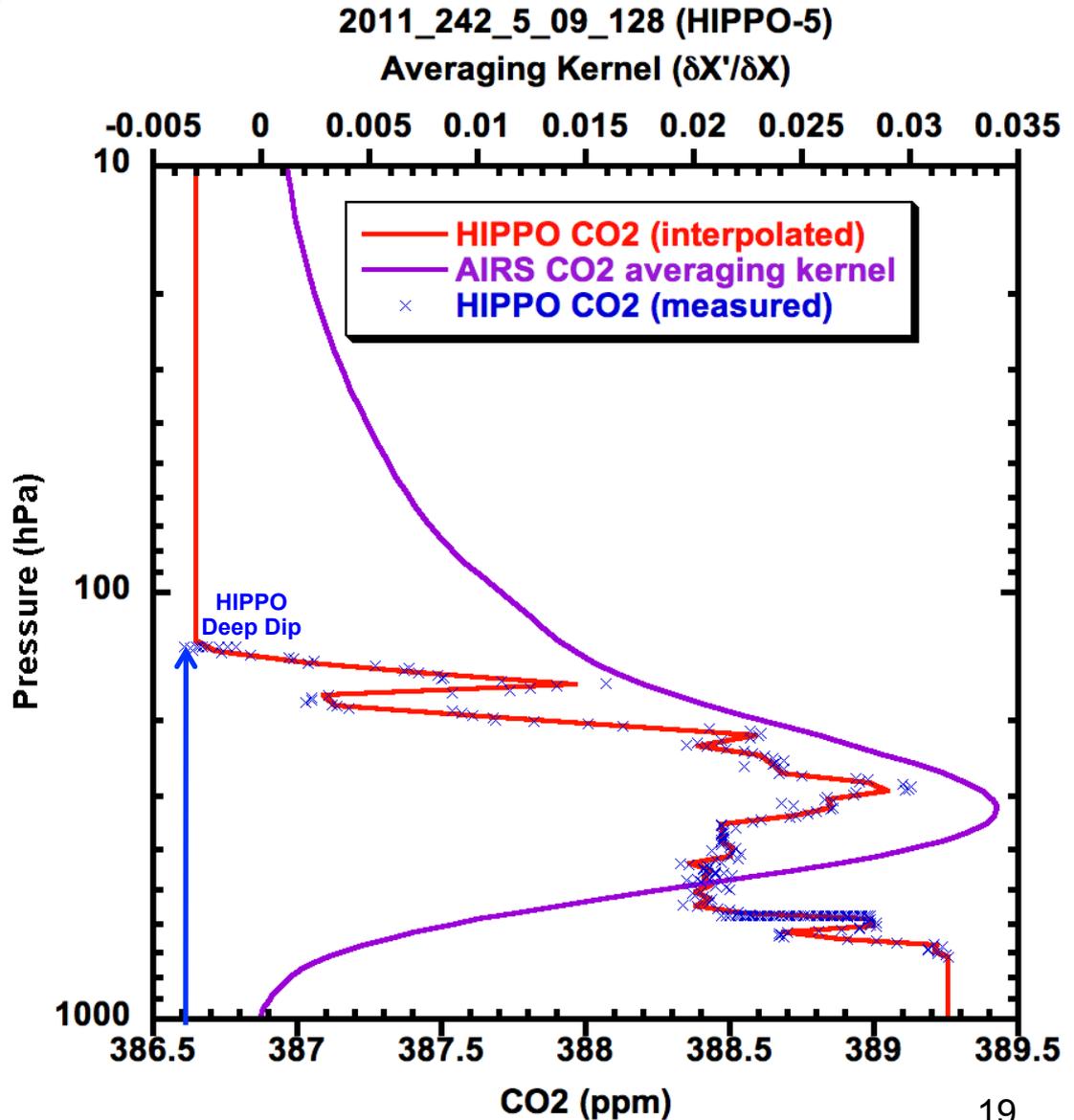
- Profiles must cover significant range of AIRS vertical sensitivity
- Convolve profiles with AIRS sensitivity function to arrive at a value to compare to collocated AIRS retrieved value
- HIPPO “Deep Dip” profiles cover sufficient vertical extent to allow validation to within 2 ppm with simple extension of profiles to higher altitudes



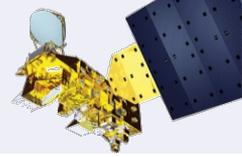


## Example

- Interpolate HIPPO measurements over range of aircraft altitude to AIRS support pressure levels
- Extend highest altitude point
- Extend lowest altitude point
- Convolve profile with AIRS sensitivity function to arrive at a value to compare to collocated AIRS retrieved value
- Extension of HIPPO “Deep Dip” profile at high altitude results in overestimation error of less than 2 ppm
- More realistic extension including stratospheric fall-off in development



# V5 AIRS/AMSU and AIRS-Only vs HIPPO

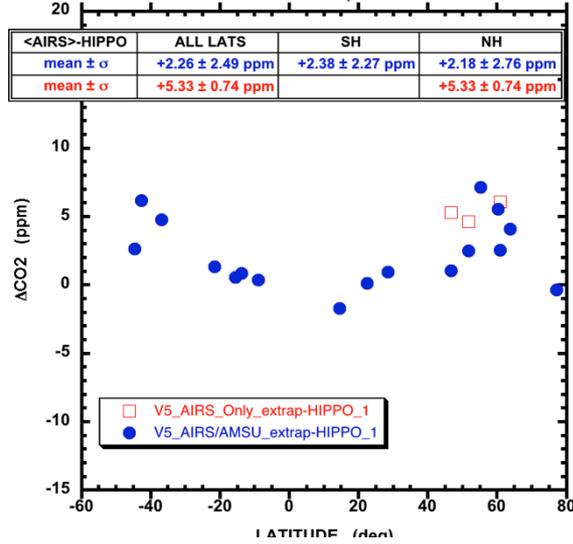


**V5 AIRS/AMSU - HIPPO-1**

**V5 AIRS-Only - HIPPO-1**

9 Jan - 22 Jan, 2009

Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km

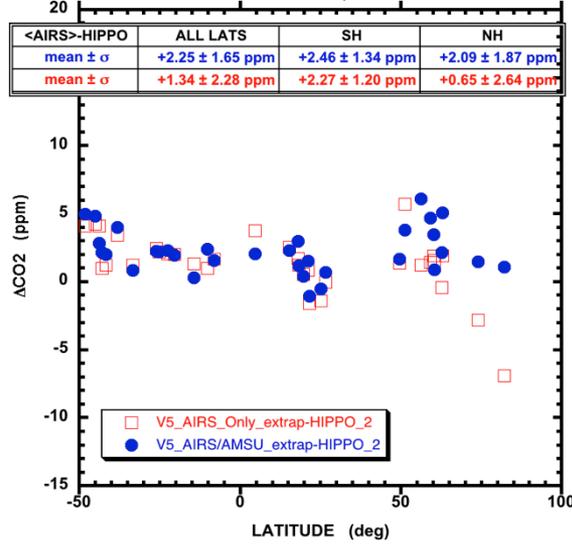


**V5 AIRS/AMSU - HIPPO-2**

**V5 AIRS-Only - HIPPO-2**

30 Oct - 22 Nov, 2009

Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km

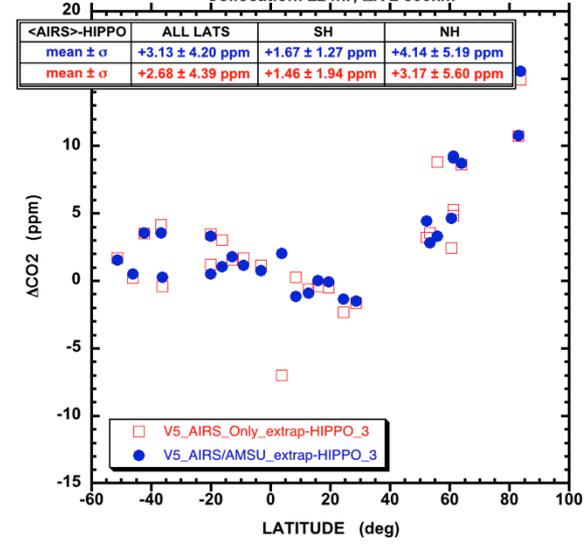


**V5 AIRS/AMSU - HIPPO-3**

**V5 AIRS-Only - HIPPO-3**

23 Mar - 16 Apr, 2010

Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km

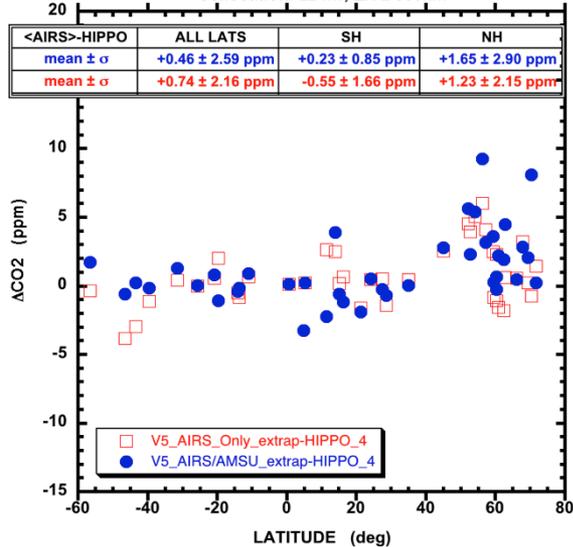


**V5 AIRS/AMSU - HIPPO-4**

**V5 AIRS-Only - HIPPO-4**

13 Jun - 30 Jun, 2011

Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km

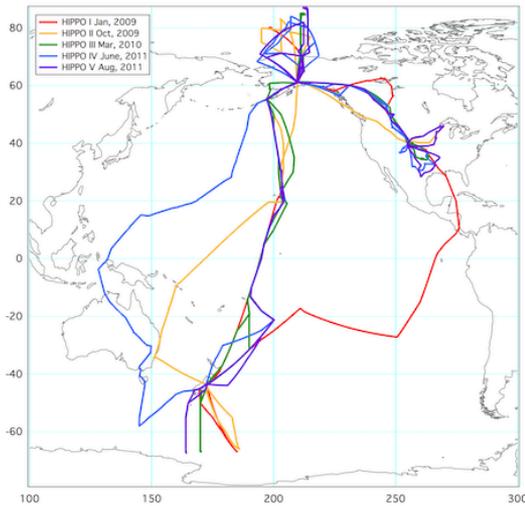
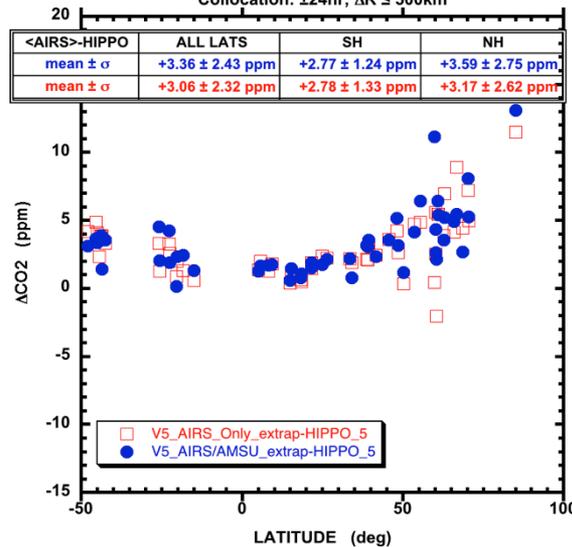


**V5 AIRS/AMSU - HIPPO-5**

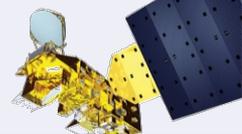
**V5 AIRS-Only - HIPPO-5**

8 Aug - 10 Sept, 2011

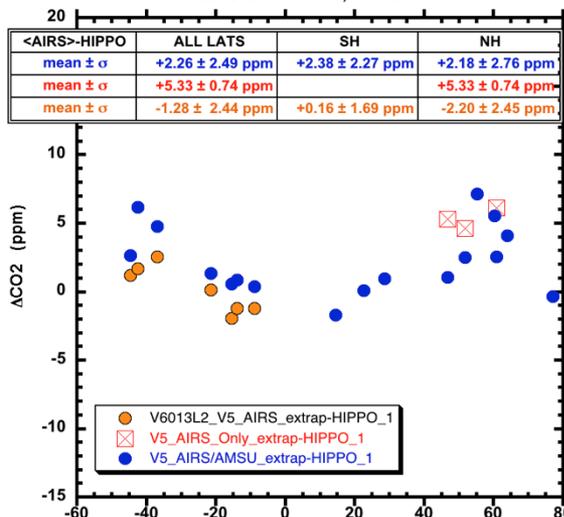
Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km



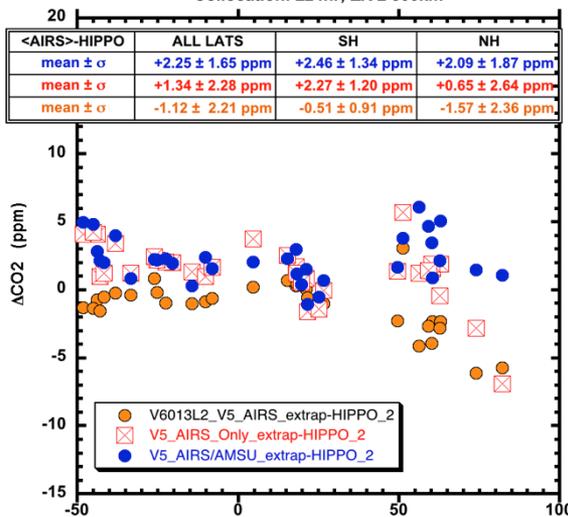
# Proto-V6 AIRS and V5 AIRS vs HIPPO



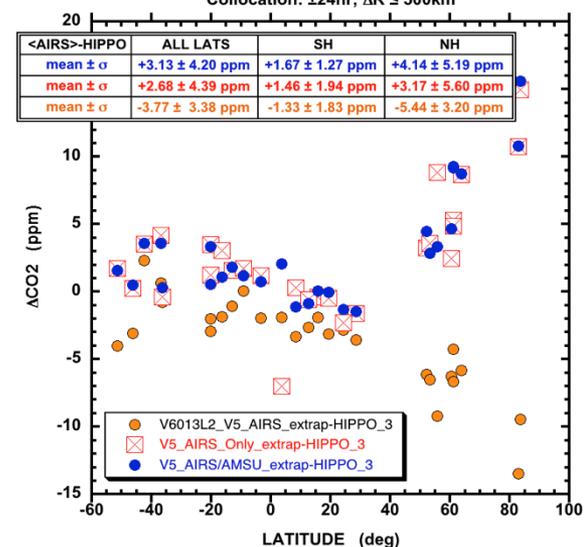
V5 AIRS/AMSU - HIPPO-1  
V5 AIRS-Only - HIPPO-1  
V6013L2\_V5\_AIRS/AMSU - HIPPO-1  
9 Jan - 22 Jan, 2009  
Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km



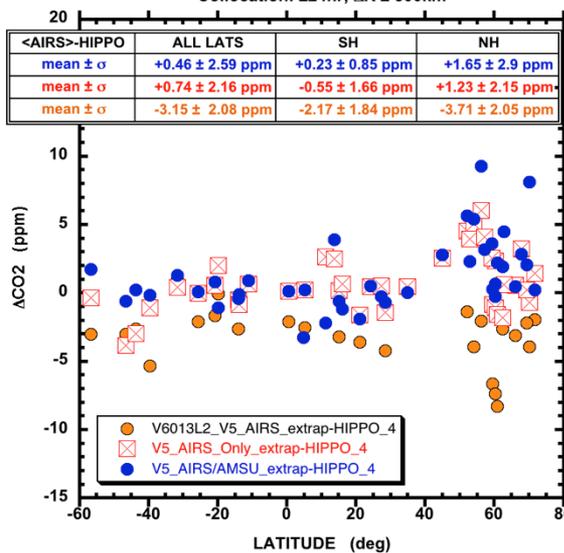
V5 AIRS/AMSU - HIPPO-2  
V5 AIRS-Only - HIPPO-2  
V6013L2\_V5\_AIRS/AMSU - HIPPO-2  
30 Oct - 22 Nov, 2009  
Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km



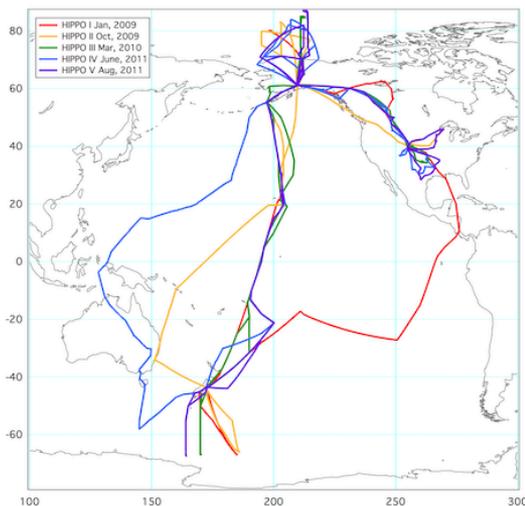
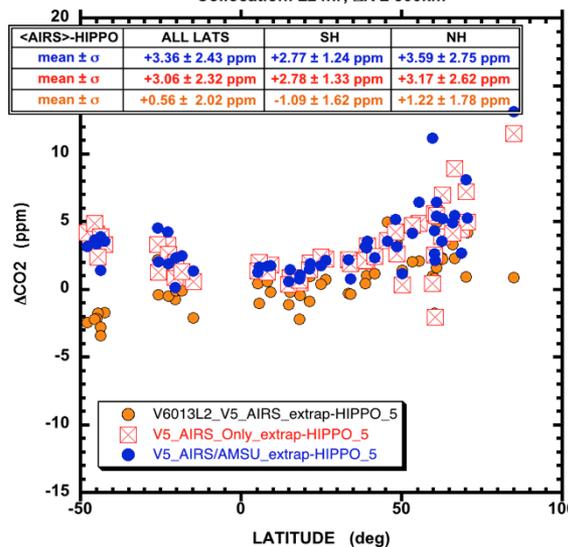
V5 AIRS/AMSU - HIPPO-3  
V5 AIRS-Only - HIPPO-3  
V6013L2\_V5\_AIRS/AMSU - HIPPO-3  
23 Mar - 16 Apr, 2010  
Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km

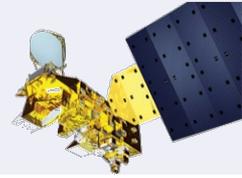


V5 AIRS/AMSU - HIPPO-4  
V5 AIRS-Only - HIPPO-4  
V6013L2\_V5\_AIRS/AMSU - HIPPO-4  
13 Jun - 30 Jun, 2011  
Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km



V5 AIRS/AMSU - HIPPO-5  
V5 AIRS-Only - HIPPO-5  
V6013L2\_V5\_AIRS/AMSU - HIPPO-5  
8 Aug - 10 Sept, 2011  
Collocation:  $\pm 24$ hr;  $\Delta R \leq 500$ km





- AIRS Mid-Tropospheric CO<sub>2</sub>
  - 11 year record provides global distribution
    - Day/Night over Ocean/Land
  - Exhibits features due to major atmospheric processes
    - Window on transport of trace gases
  - Provides some sensitivity to surface flux
    - High Correlation of CO<sub>2</sub> and GPP for July in NH Boreal Forests
    - Sensitive to surface flux over tropical region complements GOSAT (Junji Liu presentation on Monday)
- V5 AIRS-Only retrievals consistent with V5 AIRS/AMSU retrievals
  - Assures continuity of V5 record during V6 development
- V6 prototype testing under way
  - Optimizing QA
  - Investigating high-latitude difference between V5 and V6
    - Validation with properly extended HIPPO profiles